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**LANDFORMS OF THE CONTERMINOUS UNITED STATES—  
A DIGITAL SHADED-RELIEF PORTRAYAL**

by

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# **LANDFORMS OF THE CONTERMINOUS UNITED STATES— A DIGITAL SHADED-RELIEF PORTRAYAL**

**By Gail P. Thelin and Richard J. Pike**

## **VISUALIZING THE LANDSCAPE**

Realistic portrayal and mapping of topographic form is a centuries-old problem: to trick the eye into perceiving a two-dimensional graphic as a three-dimensional landscape. All traditional solutions, including those partly implemented by machine, have been artistic (Imhof, 1965; Horn, 1981). Among the cartographic devices invented by illustrators to supply the necessary visual depth cues are hachuring, hypsographic (elevation) tinting, contour density, parallel-profile density, pictorial relief, and shaded relief (Harrison, 1969; Alpha and Winter, 1971; Gilman, 1981). The latter two manual techniques have been particularly successful. Pictorial relief, which symbolizes topography by stylized morphologic types, was most fully developed in the 50 landform classes of Raisz (1931, 1939). Shaded relief, or hill shading, shows topography by the intensity of shadows cast by a light source (Imhof, 1965; Yoeli, 1965). First drafted by pencil, pen, or brush, shaded relief also has been executed by airbrush, dark-plate, and photography of raised-relief models. However, topographic detail is much too complex to be mapped both accurately and economically over large areas by any of these means.

Fast computers, analytical software, digital data, and graphic input/output devices have converged over the last three decades to largely mechanize the craft of map making (for example, Burrough, 1987, p. 4-6). This digital revolution has, among its many accomplishments, also solved the problem of mapping topographic form. Machine visualization now frees terrain portrayal from long-standing limitations (Kennie and McLaren, 1988). Topography need no longer be mapped symbolically, by discrete hand-drawn morphologic types (Raisz, 1931), or subjectively, by manual shading (Imhof, 1965). Where the necessary information is available in digital format, the computer can represent landforms as they actually are--within limitations of the source data--and portray terrain in the infinite variety of form that constitutes the true landscape. No longer is it entirely correct that maps of landforms are "drawn by men and not turned out automatically by machines" (Wright, 1942).

Yet despite the digital revolution in cartography, much truth remains in Wright's (1942) admonition that map makers are humans, not machines. Design and production of landform maps will be increasingly automated and sophisticated (Kennie and McLaren, 1988; , 1989), but the conception of a map (Yoeli, 1965, 1967) is fundamentally an intellectual rather than a mechanical process. Moreover, the portrayal of topography using digital data only now is passing from an experimental to a production technique. The many steps to a machine-made map of landforms--from data formatting, editing and processing, through image generation and correction, to preparation of a reproducible master--may be streamlined, if not wholly automated, but the sequence can involve much nonroutine trial-and-error. Such maps of topography, however technical in execution, will continue to remain the constructs of human vision and judgment.

### **MACHINE IMAGES OF TOPOGRAPHY**

Digital image-processing and computer graphics have mechanized much of the art of landform representation by combining the two most effective traditional techniques, pictorial relief and hill shading (Yoeli, 1967; Batson and others, 1975). The resulting image is a shaded pictorial-relief (physiographic) map in vertical perspective. Although automated shaded-relief maps can look deceptively like satellite pictures, they are not acquired directly by Earth-orbiting spacecraft, nor are the data from which they are made. The images are computed from a large array of closely spaced terrain heights, usually in grid-cell (raster) format, called a digital elevation (or terrain) model (DEM/DTM) (Miller and LaFlamme, 1958). Most DEMs still are made from conventional topographic maps (for details of a recent example see Hall and others, 1990).

Shaded relief is a complex derivative of terrain height (fig. 1). Digitally shaded maps of topography resemble cloud-free black-and-white aerial photographs taken at a low Sun angle, but actually they are large north-south/east-west arrays of minute gray squares (Yoeli, 1965). Each square, or picture element (pixel), represents a theoretical reflected-light intensity (brightness value) computed from a mathematical relation between ground slope, Sun position, and location of the observer (ground slope is estimated from the corresponding point and its neighbors in a DEM). The lightest and darkest tones in the image show the steepest areas; intermediate grey tones are gentle terrain. The visually most pleasing image is obtained by experimenting with location of the simulated Sun (conventionally from the northwest at 45° above the

horizon) and illumination intensity. For more information, see the section entitled "Technical Details" at the end of the discussion.

Mechanization of the shaded-relief technique for DEMs, pioneered by Yoeli (1967), has been widely applied. Batson and others (1975) made the first synoptic shaded-relief images for parts of the western United States (at 1:500,000 scale), and Arvidson and others (1982) published the first image of the conterminous United States (albeit at 1:30,000,000). These were followed by small shaded-relief maps of Australia (Moore and Simpson, 1982) and South Africa (Lamb and others, 1987), the first shaded-relief map of the Earth (Heirtzler, 1985), and a large map of Sweden (Elvhage and Lidman-Bergstrom, 1987). Among the latest synoptic images are those for the southwestern United States at 1:1,000,000 and 1:2,000,000 scale (Edwards and Batson, 1990a,b). Experimentation has further refined the technique (U.S. Geological Survey, 1986; Scholz and others, 1987; Ware, 1989), which has been equally effective in portraying gravity, aeromagnetic, geoid, and other geophysical data for interpretation (for example, Arvidson and others, 1982; Verhoef and others, 1989; and Christou and others, 1989).

Computer-generated maps such as this offer several advantages for the visual study of topography. Above all, these images portray landforms accurately and disclose their true complexity (at a given resolution), two properties that often are lost in small-scale sketches, diagrams, or conventional maps. Perhaps equally important, surface features can be viewed in a broad regional context. Unlike aerial photographs, image extent is limited only by size of the DEM. Digital shaded-relief maps also lack the distortion inherent in photographs and radar images. They are free of the vegetation and cultural features that mask topographic form on images from Landsat, SPOT, and other satellites (for example, National Geographic Society, 1976; U.S. Geological Survey, 1990). Stereopairs in shaded relief can be created digitally (Batson and others, 1975) and Sun position can be varied to obtain different views of the the same area (MOORE AND SIMPSON, 1982). Finally, shaded-relief images can be generated rapidly from digital files and do not require the time and artistic skill needed to prepare conventional relief maps.

## **APPLICATIONS OF DIGITAL LANDFORM MAPS**

Computer maps of elevation derivatives have many uses. Applications of relief shading include, but by no means are limited to, resource evaluation (Burrough, 1987) and the interpretation of regional and structural geology (Moore and Simpson, 1982; Lamb and others, 1987),

global tectonics (Verhoef and others, 1989), and geomorphology (Elvhage and Lidman-Bergstrom, 1987). Surface features in shaded relief can be studied by conventional techniques, including aerial photointerpretation. Automated relief-shading also provides an excellent cartographic base for mapping cultural and Earth science information at any scale commensurate with resolution of the source data: local (Mark and Aitken, 1990), regional (Edwards and Batson, 1990a,b), and global (Simkin and others, 1989). Shaded relief may be combined with such nontopographic information by machine registration with another digital file, for example, a computer-coded version of King and Beikman's (1974) map of United States geology.

Relief shading is only one of several ways to map topography by computer. Other derivatives of elevation include slope angle and aspect, slope curvature, local relief, and ridge and stream spacing and arrangement (Tobler, 1969; Evans, 1980; Dikau, 1988). Maps of these measures can be combined statistically to derive numerical fingerprints or signatures that characterize topography over large areas (Pike, 1988; Pike and Acevedo, 1988; Weibel, 1989). For example, the statistics of slope angle are contributing to the potential revision and elaboration of the Fenneman and Johnson (1946) physiographic units (fig. 2; Pike and Thelin, 1989). We expect that the regional geomorphology of the United States will be refined from these measures, as well as from the new shaded-relief map (for a Swedish example of such an analysis, see Elvhage and Lidman-Bergstrom, 1987). Also, maps of slope and other derivatives of elevation can be combined digitally with maps of rock type, soils, vegetation, climate, and demography, using Geographic Information Systems (GIS) technology (Burrough, 1987), to address practical problems, particularly those of land use and the environment (MchHarg, 1969). Such computer-intensive applications include mapping geologic hazards (Brabb, 1987; Pike, 1988), engineering geology (Wentworth and others, 1987), hydrologic analysis (Vieux, 1991), modeling air mass/terrain interactions for synoptic meteorology (Tesche and Bergstrom, 1978), and quantitative refinement of qualitative maps of the Nation's ecoregions (Gallant and others, 1989).

### **THE NEW DIGITAL SHADED-RELIEF MAP**

This map is the largest single-sheet graphic of relief forms of the United States since the classic hand-drawn oblique map of the same area by Raisz (1939). In concept and execution it most closely resembles Harrison's (1969) vertical dark-plate map of United States shaded relief, but is intrinsically more detailed and accurate than either of the above

maps. The new map clearly shows the regional terrain textures on which physiographic divisions of the United States were largely based (fig. 2; Fenneman and Johnson, 1946). It nicely complements Hammond's (1964) map depicting numerical classes of land-surface form and satellite-image color mosaics (National Geographic Society, 1976; U.S. Geological Survey, 1990), which emphasize vegetation and hydrography. The Raisz (1939) map, which is still available (from Raisz Landform Maps, P.O. Box 2254, Jamaica Plain, MA, 02130: 1-800-242-3199 (in MA, 617-868-3199), may be used to locate named surface features.

Greater detail is evident in the digital shaded-relief image than could, in all practicability, be included in synoptic portrayals of the Nation's terrain at this resolution by any manual technique. Much of the detail derives from the high density of the dataset (and the computer's ability to rapidly process so many terrain heights), but much of it simply reflects the map's size, which is more than twice as large as its closest predecessor (Harrison, 1969). The 1:3,500,000 scale also is the largest scale consistent with visual merging of pixels into a continuous smooth surface (contrast with Bleed and Flowerday, 1990, p. 4). Data resolution on the map, the length of a pixel edge, is 0.23 mm (0.8 km on the ground), slightly better than the 0.25 mm/pixel maximum value proposed by Yoeli (1965) for shaded-relief portrayal by computer.

Resolution of the shaded-relief data is slightly too coarse to rid the map completely of evidence of its computer origin. Minute pixel-sized regularities (aliasing) still are evident in the fine sawtooth edges of some narrow linear features. Other flaws in our map are discussed below under the section entitled "Source Data". All of these are minor compared to the information conveyed by the map and do not detract unduly from the portrayal.

### SOME LANDFORMS OF INTEREST

The new map shows geomorphic and tectonic phenomena of the United States in vivid detail. These features, which are so numerous that only a small sampling is included here (figs. 2-4), fall into two overall groups: landforms that are obvious or familiar and those that are less so. Treatises on regional geomorphology of the United States (Graf, 1987, is the most recent example) and certain small-scale maps of the country (for example, U.S. Geological Survey, 1990) may be consulted for further study of these and innumerable other features evident on the map.

The most obvious surface forms portrayed on this map are of three main types: large physiographic subdivisions, major contrasts in style of topographic expression, and small but unique landforms.

All the general surface characteristics listed on the standard map of United States physiographic units (fig. 2; Fenneman and Johnson, 1946) are crisply rendered here in shaded relief. Examples of distinctive morphologies include structural control of topography in the folded Appalachian Mountains, the tilted Sierra Nevada fault block, and the domical Black Hills (fig. 4, Nos. 1-3, respectively). The Basin and Range physiographic province can be viewed in its entirety in detail previously unavailable at this scale. Both the great roughness and structural complexity of the Rocky Mountains system (fig. 2, Nos. 16-19) and Pacific Border province (No. 24) are portrayed convincingly by mechanized shaded relief.

Also conspicuous on the map are two important contrasts in overall geomorphic style. The more fundamental is that between the static eastern United States, a passive continental margin, and the active West, where the North American and Pacific crustal plates are colliding. These differing tectonic regimes also are reflected in the Country's coastline: the Atlantic and Gulf coasts, which developed on a gently sloping continental shelf, differ markedly in detail from the Pacific coast, where the offshore profile is steep. The relief and fluvial texture of the Driftless section in southwestern Wisconsin (fig. 2, No. 12c and fig. 4, No. 4) contrast vividly on the new map with the muted drift-mantled terrain that surrounds it. This area exemplifies the many glacial/nonglacial differences in terrain across much of the northern United States (fig. 4, No. 5).

Only slightly less obvious on our map are certain small features, if familiar to the viewer or unique in morphology. Among those that can be located quickly are basalt ridges in the Connecticut River valley (fig. 4, No. 6), glacial moraines on Long Island and western Cape Cod (No. 7), two large caldera-forming volcanoes--Valles (No. 8) and Crater Lake (No. 9), and the approximate trace of the San Andreas fault zone (No. 10). Some small landforms of widespread interest, however, are found only if one knows the shape of the feature and just where to look. Most viewers seeking Mount St. Helens (No. 11), for example, will require a locational aid (such as Raisz, 1939, or a good atlas) to find it among the many other peaks in the Cascade Range.

A major strength of the new map is not so much its expression of the obvious, but rather its depiction of features that are subtler or less familiar. We have chosen examples from four main types: landforms of the central United States, low-relief landforms elsewhere, families of linear features that may have major significance, and minor linear trends.

The new image is unique in its clear portrayal of topography in the United States midcontinent. This large area contains landforms that are fully as interesting as the mountain chains to the east and west. Among these low-relief features are the broad extent and extraordinary flatness of the Mississippi Alluvial Plain (fig. 2, No. 3e) and the Llano Estacado of western Texas and eastern New Mexico (fig. 4, No. 12); the fine-grained hummocky texture of the Nebraska Sand Hills (fig. 4, No. 13)--the largest sand-dune area in the Western Hemisphere (Bleed and Flowerday, 1990); Crowley's Ridge (No. 14)--a late Pleistocene erosional remnant in the Mississippi River Embayment (near the epicenter of the 1811-1812 New Madrid earthquakes); and the many occurrences of unusual drainage patterns arising from advance and retreat of the last ice cap (No. 15).

Perhaps most striking of all midcontinent landforms shown on the map is the Coteau des Prairies (fig. 4, No. 16), "a flatiron-shaped plateau some 200 miles (300 km) long, pointing north" (Flint, 1955). The ice-scoured lowlands that flank the Coteau des Prairies in eastern South Dakota and neighboring States formerly were occupied by the James and Des Moines lobes of the last ice sheet. These lowlands, which drained melting ice and various ice-dammed lakes during deglaciation of the region, bear a remarkable resemblance to one type of channel observed on spacecraft images of the planet Mars (Kehew and Lord, 1986).

Subtle or only locally important landforms elsewhere in the country also are well shown on the new map. These include the enigmatic Sutter Buttes volcano (fig. 4, No. 17) in California's Sacramento Valley, the many low volcanic shields on the Snake River plain (No. 18), Trail Ridge (No. 19, an important domestic source of the minerals ilmenite and zircon) and other long north-trending sandy ridges in Florida, the Clinton-Newbury (or Essex) fault zone (No. 20) and its extensions in eastern Massachusetts and Connecticut (a minor crustal suture), and the overall north-northeast-trending texture of the Appalachian Mountains in northern New England (No. 21).

Groups of throughgoing linear features, which may mark significant events in the tectonic evolution of the North American continent, are among the most remarkable patterns on this map. Some of them, such as



the west-trending lineament in the Rocky Mountains (fig. 4, No. 22), the east-northeast trend that aligns with the Garlock fault zone (No. 23), and an east-northeast trend (No. 24) that includes parts of the Gila and Salt Rivers (Arizona) and the Canada River (Texas), are perhaps seen clearly for the first time here. The latter two trends parallel the Murray fracture zone and other inactive transform faults (not shown) of the eastern Pacific plate. Major elongate features that are more familiar from previous study include the Olympic-Wallowa lineament (No. 25) and the many, possibly related, subparallel northwest trends in the Rocky Mountains (No. 26) to the southeast of it. Some of the large lineaments on this map may include currently active faults .

Minor trends in some parts of the map also may reflect neotectonic activity. However, interpreting such trends requires restraint, as it does for aerial photographs and radar images in certain types of topography (Howard and Larsen, 1972; Yamaguchi, 1985) . On this map, the 300° lighting slightly exaggerates 325°-striking ridges and valleys, especially those that produce the remarkable north-northwest grain on the Great Plains (fig. 4, No. 27), which seems to be real nonetheless (unidirectional lighting at a low elevation angle may artificially enhance small linear features in preferential directions, typically 20° to 35° from that of the light source, while suppressing them in others, parallel to the light source). The north-northwest trend may mark a regional fracture pattern imprinted on the middle one-third of the Country, reflecting its coupling a mobile West to a more stable East. It suggests indeed "that the effect of plate collisions may extend across vast distances upon the surface of a continental plate" (Beckinsale and Beckinsale, 1989).

## FUTURE PROSPECTS

Our shaded-relief image of the United States is not necessarily the ultimate digital portrait of the Nation. Like any reconnaissance geologic map, or for that matter a good scientific hypothesis, it is an ongoing experiment (Yoeli, 1965). Because this map reflects a still-evolving technology, various improvements are under consideration. Foremost among these are restoration of digital elevations for southern Canada and northern Mexico (see following section entitled "Technical Details"; Arvidson and others, 1982), inclusion of Alaska and Hawaii, eliminating or reducing errors in the dataset through further editing and edge-matching of data blocks, and more hydrography. The visual perception of elevation could be enhanced through the use of color (Harrison, 1969; Lamb and others, 1987).

Changes in the shaded-relief calculation might address some remaining shortcomings of the map, particularly tonal imbalance between steep and gentle terrain. Detail in very mountainous areas is obscured because the steepest slopes are too dark or too light. We have found that the desired balance in tone can not be achieved simply by transforming all elevation or slope values to logarithms or square roots, and then computing brightnesses from the transformed values (U.S. Geological Survey, 1986). The solution is likely to be more complex and may require incorporating special algorithms, called local operators, within the computer software to tailor reflectance values to specific conditions of elevation and slope (Brassel, 1974).

Lastly, the map could be best improved simply by increasing scale, data density, and image resolution (for example, Edwards and Batson, 1990a,b). Such a map would provide the detail needed for a more effective relief portrait in many parts of the United States (for example, Bleed and Flowerday, 1990, p. 4). An improved map of the entire country, probably at 1:1,000,000 scale and necessarily in several sheets, would require a cleaned-up file of all the original digitized elevations and an image resolution of about 0.1 mm/pixel (130 m on the ground). Multiple editions of this map at several Sun azimuth and elevation settings (Moore and Simpson, 1982) would further exploit the research potential of digital shaded-relief by accentuating terrain features that follow all the different trends in the Country's landscape.

## **TECHNICAL DETAILS**

### **IMAGE PROCESSING AND HILL SHADING**

Our map was made by digital image-processing, a technical specialty related to the broader fields of computer graphics and machine vision (Dawson, 1987; Kennie and McLaren, 1988). The technology includes the many spatially based operations first brought together and developed systematically to manipulate Ranger, Mariner, Landsat, and other images that are reassembled from spacecraft telemetry in a raster or scan-line arrangement of square-grid elements (Nathan, 1966; Castleman, 1979; Sheldon, 1987). These computer procedures have been successfully transferred to landform analysis from remote-sensing applications by substituting terrain heights or sea-floor depths for the customary values of electromagnetic radiation obtained from satellites and stored in digital arrays of pixels (Batson and others, 1975).

It is the 1 pixel=1 elevation equivalence that enables image-processing technology to so efficiently map elevation matrices and their derivatives over large areas. Recent examples are given by Arvidson and others (1982), Simpson and others (1986), and Verhoef and others (1989). Slope angle and other quantitative measures of surface form can be rapidly calculated, compared, and combined for display as shaded-relief and color images or stored as digital files for further study of topography and registration with nontopographic datasets (Batson and others, 1975; Moore and Mark, 1986; Pike and Acevedo, 1988; Pike and Thelin, 1989).

The image-processing tool applied here, relief shading, is more formally termed analytical hill shading. Although well known as an artistic and manual technique (Imhof, 1965), it was impractical for large areas until Yoeli (1965) developed a modern analytical version for square-grid matrices of terrain elevations and then automated it for the computer (Yoeli, 1967). Analytical hill shading maps topographic form through variations in mathematically determined intensity of reflected light ( $I$ ) at each elevation/ pixel located on the ground (fig. 5). This relation, known as the photometric function, has many variants (Brassel, 1974; Batson and others, 1975; Horn, 1981; Christou and others, 1989). The simplest case is the cosine law of Lambert

$$I = k_d \cos(i) \quad (1)$$

where  $i$  is the angle between the incident light (the Sun) and a vector normal to the sloping ground, and  $k_d$  is a coefficient describing reflectivity of the surface material (here, a perfect diffuser of incident light; Greenberg, 1989). Position of the viewer is directly overhead. Ground slope may be estimated from a DEM in several ways, using three to nine adjacent height values (Ware, 1989; Mark and Aitken, 1990). Repetition of these calculations pixel-by-pixel over a large DEM yields a reflectance map, a continuous X,Y array of brightness values (Horn, 1981).

Many refinements to the basic approach itself can improve relief shading without having to add data from other sources (such as Landsat). Besides direct illumination, reflectance maps generally include some ambient light, which strikes and reflects from a surface equally in all directions, to improve appearance of the final image (Greenberg, 1989; Christou and others, 1989). Shadows cast by steep terrain also can be incorporated into the calculation (Ware, 1989), and even atmospheric effects can be simulated (Brassel, 1974). Finally, advanced techniques of computer graphics used in some industries to digitally depict virtually any object with photorealistic quality (Whitted, 1982; Greenberg, 1989)

conceivably could be adapted to take shaded-relief portrayal to even higher levels of realism (Kennie and McLaren, 1988).

## SOURCE DATA

The terrain heights from which our map was made do not owe their origin to remote sensing by Landsat or other spacecraft. Rather, the data result from the machine sampling--initially by contour-tracing, later by drum-scanning--of available contour maps, some of which were first compiled as early as 1947. These measurements have a complex history that spans a quarter of a century, starting with the Defense Mapping Agency Topographic Center's (DMATC) creation of a nationwide set of gridded elevations in 1964 to 1972.

DMATC digitized and labeled contour lines, and later spot heights and stream and ridge lines, on hundreds of 1:250,000-scale (1° by 2°) topographic sheets covering the United States and parts of Canada and Mexico. Digitizing these maps by semi-automated methods at 0.01 in. (0.25 mm) resolution, 3 arcseconds or about 200 ft (63 m) on the ground, accounted for 1/6 of the elevations (Mays, 1966). The remaining 5/6 of the data were interpolated between digitized contours by computer (Noma, 1966). The entire DEM, containing more than 2 billion elevations arranged in a square grid of 3-arc-second resolution, has been available since 1974 in over 900 10 by 10 blocks from the U.S. Geological Survey (USGS), currently the Earth Science Information Center (Rm. 1C107, 507 USGS National Center, Reston VA 22092: 1-800-860-6045).

The original DMATC data were later resampled (thinned) and averaged down (see Godson, 1981, for some details) to the more manageable file used here and by Godson (1981) and by Arvidson and others (1982). The resulting 12 million elevations are spaced 30 arc-seconds apart, nominally 0.805 km on the ground, north-south and east-west. The actual array (6,046 by 3,750) processed to make our map includes null (black) background values lying between the national boundary and the map border and thus is much larger. Although the initial DMATC data were read or interpolated to the nearest foot (Mays, 1966), the elevations were later rounded to 10 m (map contour intervals were coarse: 100 ft or more). Accordingly, vertical accuracy of the data for this image may be no more than 30 m in smooth areas and 50 m in rough terrain. The 30-arc-second DEM is available from the National Oceanic and Atmospheric Administration's National Geophysical Data Center (Code E/GC1, 325 Broadway, Boulder CO 80303: 303 497-6128).

Errors in both the 3-arc-second and the 30-arc-second datasets, in addition to those inherent in the source maps, account for visible flaws in the map (Batson and others, 1975). Most of these errors are systematic. The flattened hills and the fine-scale rectilinear and stair-step textures on the map arise from round-off error and also from inaccurate interpolation between widely spaced contours, the result of too large a contour interval and a fast but suboptimal algorithm dictated by the slow computers available 25 years ago (Noma, 1966). Faint, widely spaced north-south and east-west lines mark defective splices between 1° data blocks. Less systematic errors include additional flattened hilltops, unexplained textures, and randomly located zero or unduly high elevations arising from unknown causes.

## COMPUTATION AND PRODUCTION

We created this map by processing all of the 30-arc-second height data through proprietary software, the Interactive Digital Image Manipulation System (IDIMS; Electromagnetic Systems Laboratory, Inc., 1983), installed on a DEC VAX 11/780 computer. After registering the location of each of the 12 million elevations to an Albers Equal-Area Conic projection (standard parallels at 29.5° N and 45.5° N; central meridian at 96.0° W, and latitude of projection's origin at 23.0° N), we produced a new grid of 0.805-km-resolution pixels from bilinear resampling. Topography beyond the national boundary, in two strips across southern Canada and northern Mexico, was excluded from the dataset by a 1:2,000,000-scale United States outline obtained from a USGS digital line graph.

The SUNSHADE routine within IDIMS computed strike and dip angles of terrain slope, by algebraic manipulation of the four elevations immediately north, south, east, and west of each sample point in the DEM (fig. 6), and from them assigned brightness values ranging from 0 (deeply shadowed areas) to 255 (fully illuminated surfaces) to all 12 million pixels (these calculations took about 17 minutes of CPU time on the VAX 11/780). The algorithm (Electromagnetic Systems Laboratory, Inc., 1983) is built around a much-modified Lambertian photometric function (Horn, 1981) that uses diffuse scattering to simulate the effects of solar illumination,

$$I = k_d [L \cos(i)] + A \quad (2)$$

where  $L$  is a scaling factor for the intensity of illumination, and  $A$  is an additive ambient light factor (the remaining terms are defined for

equation 1, above). The calculation does not provide for cast shadows (Ware, 1989). Various parameters to the SUNSHADE routine control image contrast and thus final appearance of the shaded relief. We found that the following values gave the most crisp and visually appealing portrayal overall: vertical exaggeration, 2X; Sun azimuth, 300°; Sun elevation, 25°; scaling factor for the intensity transformation, 1.2 units; ambient-light factor, 0.7 units.

Errors in the DEM were located from both statistical analysis of the elevations (Pike and Thelin, 1989) and visual identification of aberrant patterns in the image. We repaired some of the most visible artifacts by editing flawed portions of the map and changing brightness values on a pixel-by-pixel basis, using a Hewlett-Packard 9000 Series-360 Turbo workstation. To retain maximum local detail in the map, we did not attempt to correct or change any of the erroneous elevations globally, by applying a digital filter to the entire DEM.

Preparation of the final image required several steps. To increase tonal contrast in smooth topography and diminish it in areas of high relief, we remapped the intensity values output from SUNSHADE, using a piecewise linear transformation, to new values based on breakpoints that we defined in the original shaded-relief image from a histogram of its brightnesses. Over 70 of the largest (over 20 mi<sup>2</sup> in area) natural lakes--plus the coastline and national boundary-- were obtained from an analog 1:2,500,000-scale United States map and used to mask the shaded-relief file photomechanically as a nondigital overlay (shown in black on the map).

The printed map was created from three separate negatives made on a Scitex Response-300 computerized cartographic system. To obtain the desired tonal contrast and balance, each negative emphasizes a different range of brightness. The full range of light intensities in each negative was computed for the 12 million pixels to generate a printing screen of 175 lines/inch (an image resolution of 0.15 mm) on the Scitex system's laser drum plotter. The final map was printed in black ink, from five lithographic plates—three made from the shaded-relief negatives and two made from the nondigital overlay to insure a deep black for the background.

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## FIGURE CAPTIONS

Figure 1. Basic steps in the generation of a shaded-relief map (modified from Brassel, 1974; and Horn, 1981).

Figure 2. Physiographic divisions (heavy lines), provinces (numbers), and sections (letters) of the conterminous United States (from Fenneman and Johnson, 1946).

**EXPLANATION**

LAURENTIAN UPLAND

- 1 Superior Upland

ATLANTIC PLAIN

- 2 Continental Shelf (not on map)

- 3 Coastal Plain

- a Embayed section
- b Sea Island section
- c Floridian section
- d East Gulf Coastal Plain
- e Mississippi Alluvial Plain
- f West Gulf Coastal Plain

APPALACHIAN HIGHLANDS

- 4 Piedmont province

- a Piedmont Upland
- b Piedmont Lowlands

- 5 Blue Ridge province

- a Northern section
- b Southern section

- 6 Valley and Ridge province

- a Tennessee section
- b Middle section c Hudson Valley

- 7 St. Lawrence Valley

- a Champlain section
- b Northern section (not on map)

- 8 Appalachian Plateaus province

- a Mohawk section
- b Catskill section
- c Southern New York section
- d Allegheny Mountain section
- e Kanawha section
- f Cumberland Plateau section
- g Cumberland Mountain section

- 9 New England Province

- a Seaboard Lowland section
- b New England Upland section
- c White Mountain section
- d Green Mountain section
- e Taconic section

- 10 Adirondack province

## INTERIOR PLAINS

- 11 Interior Low Plateaus
  - a Highland Rim section
  - b Lexington Plain
  - c Nashville Basin
- 12 Central Lowland
  - a Eastern Lake section
  - b Western Lake section
  - c Wisconsin Driftless section
  - d Till Plains
  - e Dissected Till Plains
  - f Osage Plains
- 13 Great Plains province
  - a Missouri Plateau, glaciated
  - b Missouri Plateau, unglaciated
  - c Black Hills
  - d High Plains
  - e Plains Border
  - f Colorado Piedmont
  - g Raton section
  - h Pecos Valley
  - i Edwards Plateau
  - k Central Texas section

## INTERIOR HIGHLANDS

- 14 Ozark Plateaus
  - a Springfield-Salem plateaus
  - b Boston "Mountains"
- 15 Ouachita province
  - a Arkansas Valley
  - b Ouachita Mountains

## ROCKY MOUNTAIN SYSTEM

- 16 Southern Rocky Mountains
- 17 Wyoming Basin
- 18 Middle Rocky Mountains
- 19 Northern Rocky Mountains

## INTERMONTANE PLATEAUS

- 20 Columbia Plateau
  - a Walla Walla Plateau
  - b Blue Mountain section
  - c Payette section
  - d Snake River Plain
  - e Harney section
- 21 Colorado Plateaus

- a High Plateaus of Utah
- b Uinta Basin
- c Canyon Lands
- d Navajo section
- e Grand Canyon section
- f Datil section
- 22 Basin and Range province
  - a Great Basin
  - b Sonoran Desert
  - c Salton Trough
  - d Mexican Highland
  - e Sacramento section
- PACIFIC MOUNTAIN SYSTEM
- 23 Cascade-Sierra Mountains
  - a Northern Cascade Mountains
  - b Middle Cascade Mountains
  - c Southern Cascade Mountains
  - d Sierra Nevada
- 24 Pacific Border province
  - a Puget Trough
  - b Olympic Mountains
  - c Oregon Coast Range
  - d Klamath Mountains
  - e California Trough
  - f California Coast Ranges
  - g Los Angeles Ranges
- 25 Lower California province

Figure 3. Boundaries of the conterminous 48 United States.

Figure 4. Topographic features cited in discussion of shaded-relief map of conterminous United States.

#### **EXPLANATION**

- 1 Folded Appalachian Mountains
- 2 Sierra Nevada
- 3 Black Hills
- 4 Driftless area
- 5 Southern limit of Pleistocene continental glacial deposits
- 6 Basalt ridges in the Connecticut River valley



- 7 Terminal moraines on Long Island and Cape Cod
- 8 Valles caldera
- 9 Crater Lake caldera
- 10 San Andreas fault zone
- 11 Mount St. Helens
- 12 Llano Estacado
- 13 Nebraska Sand Hills
- 14 Crowleys Ridge
- 15 Unusual drainage patterns emplaced during Pleistocene
- 16 Coteau des Prairies
- 17 Sutter Buttes
- 18 Low volcanic shields on Snake River plain
- 19 Trail Ridge
- 20 Clinton-Newbury (Essex) fault zone
- 21 Linear texture in northern New England
- 22 West-trending lineaments in Rocky Mountains
- 23 Extension of Garlock fault zone
- 24 Linear trend following Gila, Salt, and Canada Rivers
- 25 Olympic-Wallowa lineament
- 26 Northwest trends in Rocky Mountains
- 27 North-northwest grain on High Plains

Figure 5. Geometric relation between ground slope and Sun position that is basic to reflected-intensity (brightness) calculations for shaded-relief mapping (modified from Batson and others, 1975). See equations (1) and (2) in text. Point O is center terrain height in five-point sample of the digital elevation model (see fig. 6).

Figure 6. Obtaining values of strike and dip for local terrain slope within a square-grid digital elevation model (subset of 24 heights shown here by open circles) from a five-point sample design (filled circles). Calculation for center terrain height, O (see fig. 5), averages east-west and north-south slope values defined by neighboring heights  $X_1$  and  $X_2$  and  $Y_1$  and  $Y_2$ , respectively (Electromagnetic Systems Laboratory, Inc., 1983). Sample point O is relocated at each height value throughout the DEM and the calculation repeated. Resulting values of strike and dip are used to compute the slope normal and then angle  $i$  (fig. 5) for text equations (1) and (2).